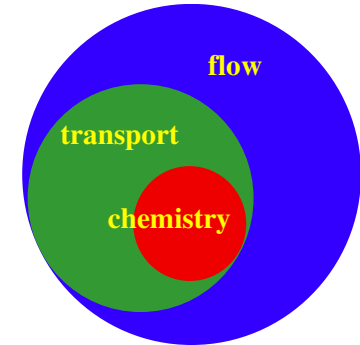


Multicomponent Geochemical Computations in TOUGHREACT

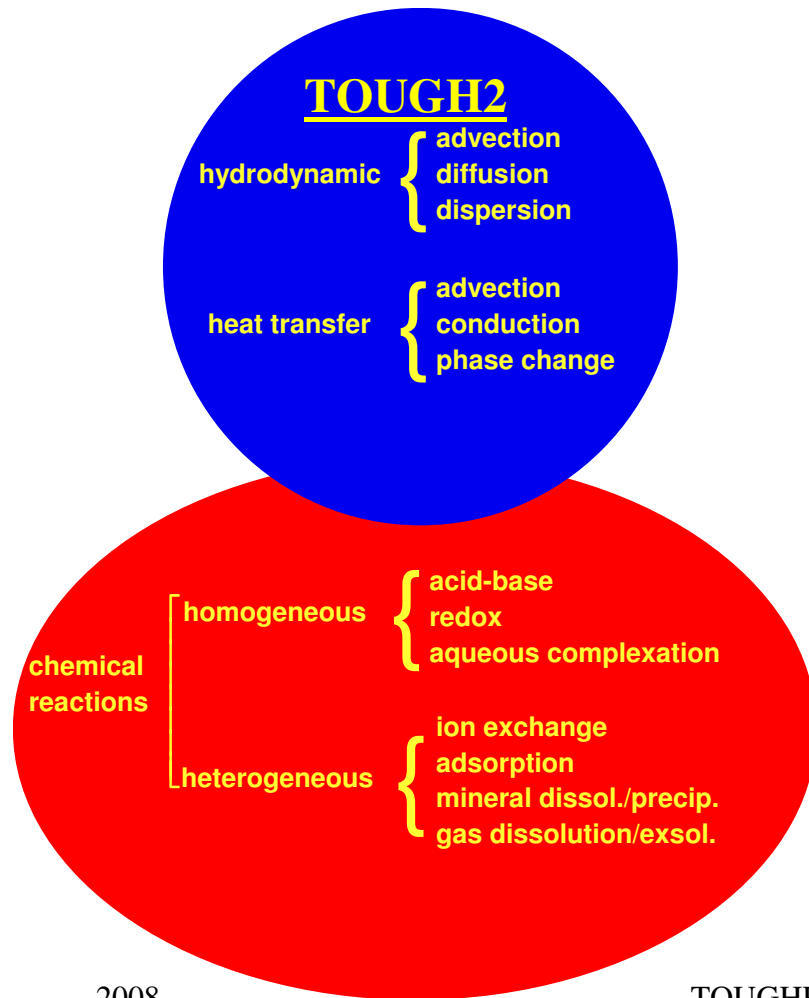
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TOUGHREACT

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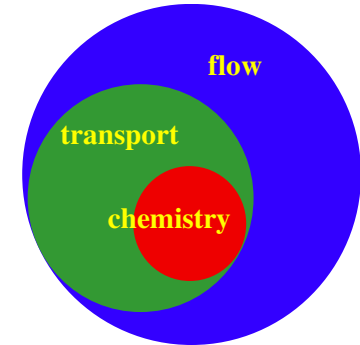
Processes



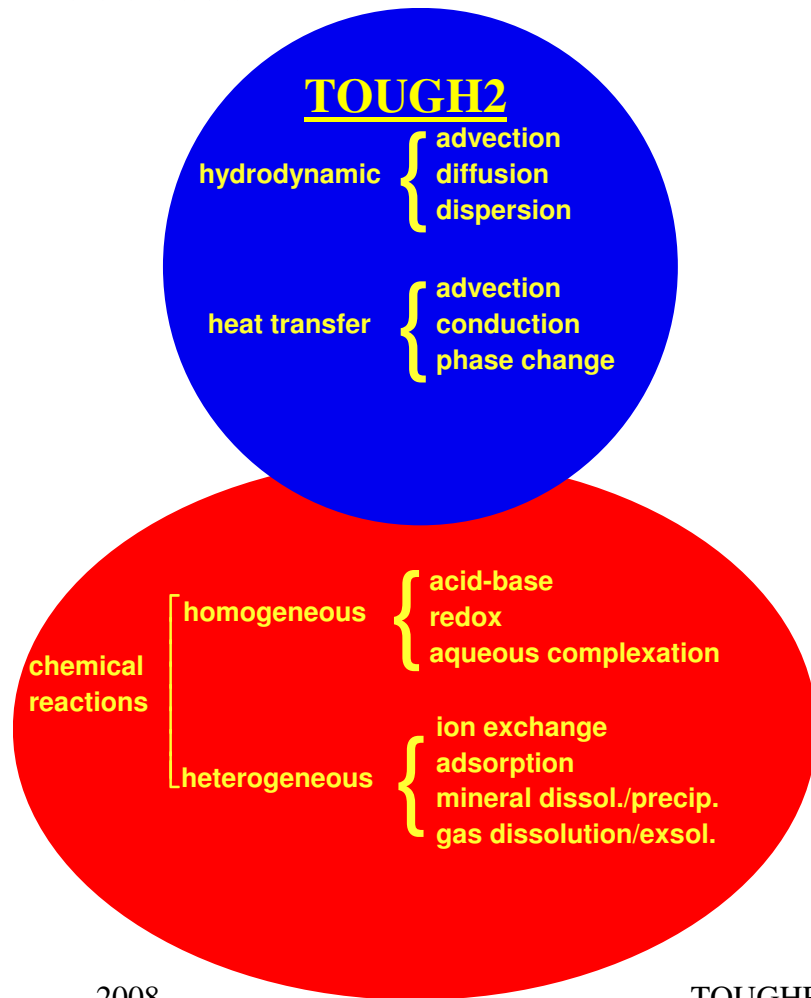
- Introduction of reactive chemistry into TOUGH2
- Reactions between gas - aqueous - solid phases, equilibrium or kinetics
- General database for minerals, aqueous and gaseous species
- Porosity and permeability change

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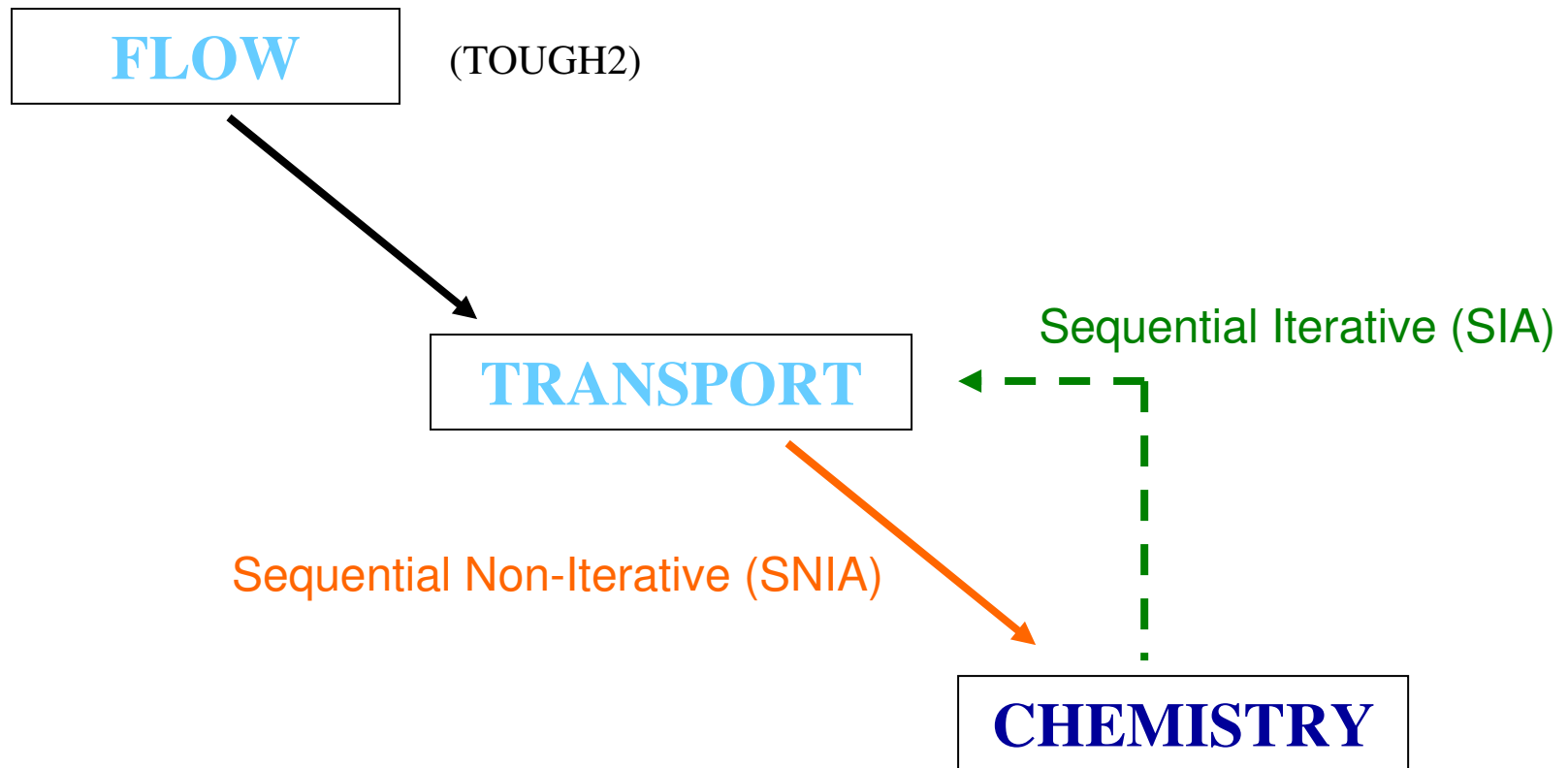
Processes



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Focus

Reactive Transport Computational Approach



Multicomponent Chemical System

- Multicomponent Reactions (water must be present)
 - Aqueous Species
 - Minerals
 - Gases
 - Exchange Species
 - Surface Complexes (v2.0 beta)
- Aqueous Speciation: equilibrium (v2.0 beta with kinetics)
- Surface Complexation/Exchange: equilibrium
- Mass Transfer:
 - Minerals: equilibrium or kinetic constraints
 - Gases: equilibrium
- External Thermodynamic Database: reaction stoichiometries, equilibrium constants, activity coefficient data, etc.

Geochemical Computations - General

- Chemical system definition
 - Temperature and Pressure
 - Total aqueous concentrations (e.g., analytical)
 - Mineral amounts (0 for potential secondary phases)
 - Gas partial pressures (optional, unsaturated medium)
- Numerical approach: mass balance/mass action
 - Primary aqueous species (actual unknowns)
 - Derived species (functions of primary species – mass action)
 - Secondary aqueous species (ion pairs, complexes)
 - Minerals
 - Gases
 - Newton-Raphson iterative procedure
 - Solve for concentrations of primary species and kg water
 - Derive all other concentrations, mineral amounts, and gas partial pressures from mass action laws involving primary species

Example Chemical System

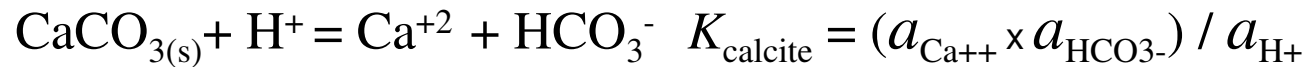
- Components of interest:
 - O, H, Na, Cl, Ca, S, C
- Primary species reflecting the components:
 - H_2O , H^+ , Cl^- , Ca^{+2} , Na^+ , SO_4^{-2} , HCO_3^-
 - Best to use primary species representing dominant species
 - H_2O and H^+ always primary species in TOUGHREACT
- Secondary aqueous species
 - OH^- , CO_3^{-2} , HSO_4^- , CaHCO_3^+ etc... (automatic selection or specified)
- Minerals (specified input)
 - Calcite (CaCO_3), Gypsum (CaSO_4) etc...
- Gases (specified input)
 - $\text{CO}_{2(g)}$, $\text{HCl}_{(g)}$ (Note, $\text{H}_2\text{O}_{(g)}$ is handled by flow EOS modules!)

Mass-Action Equations

- Expressions in terms of primary species only
- Secondary aqueous species ($a_i = \gamma_i \times m_i$)



- Minerals



- Gases ($f_i = \phi_i \times P_i$)



Mass-Balance Equations

- Total moles (M^t) in terms of molal concentrations (mol/kg_w)

- Solutes

substitute mass action (in terms of γ and m)

$$M^t_C = \{[\text{HCO}_3^-] + [\text{CO}_3^{-2}] + [\text{CaHCO}_3] \dots\} \times \text{kg}_w + \Delta n_{\text{calcite}} + n_{\text{CO}_2(\text{g})} + \dots$$

$$M^t_{\text{H}^+} = \{[\text{H}^+] - [\text{OH}^-] - [\text{CO}_3^{-2}] + [\text{HSO}_4^-] + \dots\} \times \text{kg}_w - \Delta n_{\text{calcite}} + n_{\text{CO}_2(\text{g})} + \dots$$

$$M^t_{\text{Ca}} = \{[\text{Ca}^{+2}] + [\text{CaHCO}_3^+] + \dots\} \times \text{kg}_w + \Delta n_{\text{calcite}} + \Delta n_{\text{gypsum}} + \dots$$

- Solvent (water, after Reed 1982 GCA)

$$M^t_w = \{55.505 + [\dots] + \dots\} \times \text{kg}_w + \Delta n_{\dots} - n_{\text{CO}_2(\text{g})} + \dots$$

- Minerals – incremental change Δn_m (in moles)

Kinetics: $\Delta n_m = R \times \Delta t$ $R = \text{kinetic rate}$ (positive = dissolve)

Equilibrium: solve for Δn_m by adding mass action eqn (negative = dissolve)

- Gases, at equilibrium – total moles n_i (change + initial)

From gas law: $P_i V = n_i RT$ with gas concentration = n_i/V and P_i from mass-action law (ideal, $P = f$, except with ECO2 module)

Activity Coefficients Calculations

- Charged species and water:
Extended Debye-Hückel (Helgeson et al., 1981 AJS)
 - Assumes predominant ions are Na^+ and Cl^-
 - Careful above ionic strength 1 molal for non Na-Cl dominant waters!
 - Not bad to ~ 6 m for pure NaCl solution
 - Remove $\text{NaCl}_{(\text{aq})}$ (derived species) in the database!
 - Watch! D-H a_0 values are calculated from input effective radii in database (different from more standard “b-dot” equation)
- Neutral species:
 - Optional Setchenow equation: $\log(\gamma) = (\text{const}) \times (\text{ionic strength})$, otherwise γ values are set to one
 - Dissolved gases: Drummond (1981 – unpubl. Ph.D. Penn State)
- Unreleased beta-version with Pitzer ion-interaction model for concentrated solutions (Zhang et al., 2006)

Redox Reactions

- Same approach as non-redox reactions if:
 - Unique redox couple (i.e., donor and acceptor) is defined with two primary species. For example:
 - H_2O and O_2 (preferable if dealing with mostly oxidized systems)
 - SO_4 and HS^- (preferable if dealing with mostly reduced systems)
 - Electron transfer is balanced using the specific redox couple in all reactions of the database (i.e., no free electrons!).
- Example:
 - Primary species: Fe^{+3} , H_2O , $\text{O}_{2(\text{aq})}$, etc... (with Fe^{+2} as secondary species)
$$\text{Fe}^{+2} + 0.25\text{O}_{2(\text{aq})} + \text{H}^+ = \text{Fe}^{+3} + 0.5\text{H}_2\text{O}$$
 - Primary species: Fe^{+2} , SO_4 and HS^- , etc... (with Fe^{+3} as secondary species)
$$8\text{Fe}^{+3} + \text{HS}^- + 4\text{H}_2\text{O} = 8\text{Fe}^{+2} + \text{SO}_4^{--} + 9\text{H}^+$$
- Utility provided to “switch” the redox couple in the database
- Unreleased v2.0 beta with redox disequilibrium (Xu, 2006)

Thermodynamic Data

- $\log(K)$ values for mass-action equations and D-H parameters for activity coefficients are calculated as functions of temperature (T_K , *Kelvin*) from coefficients read in external database for given reaction stoichiometries

$$a + \ln(T_K) + b + cT_K + d/T_K + e/T_K^2$$

- **Watch!!!! Crucial for confidence in results!**
 - Quality/consistency of $\log(K)$ data
 - Applicability of activity coefficient model
 - Consistency between activity coefficient model and types of secondary aqueous species and their $\log(K)$ values
- DO NOT use supplied database as black box!
- Understand the data you are using (trash in = trash out)

Kinetic Data

$$Rate = \pm k A_m \prod_i a_i^p \left[\left(\frac{Q}{K} \right)^m - 1 \right]^n$$

(e.g., Steefel and Lasaga, AJS, 1994)

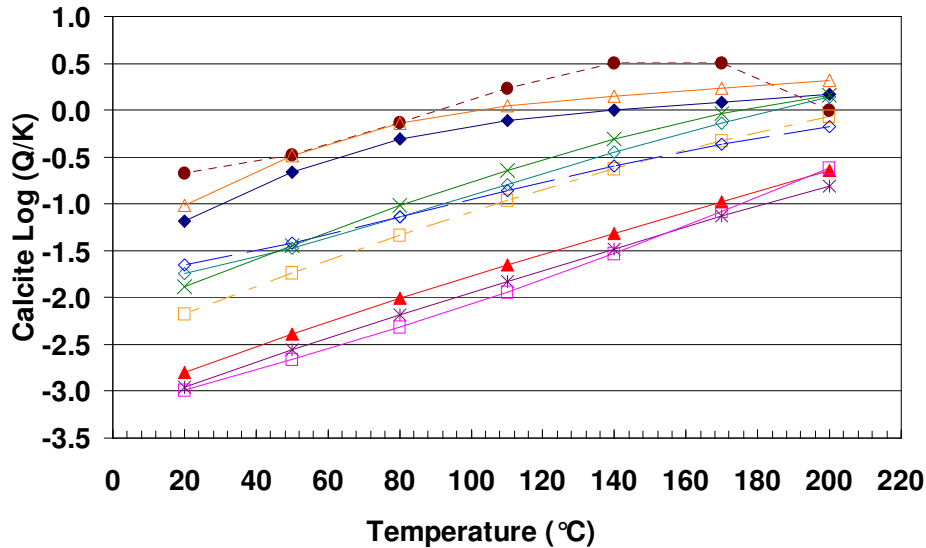
- Rate constant k
- Surface area A_m
- Equilibrium constant K

$$k = k_0 \exp \left[\frac{-E_a}{R} \left(\frac{1}{T} - \frac{1}{298.15} \right) \right]$$

Q is ion activity product, a_i are individual activities (e.g., H⁺)
Equilibrium at Q/K = 1, dissolution at Q/K < 1, precipitation at Q/K > 1

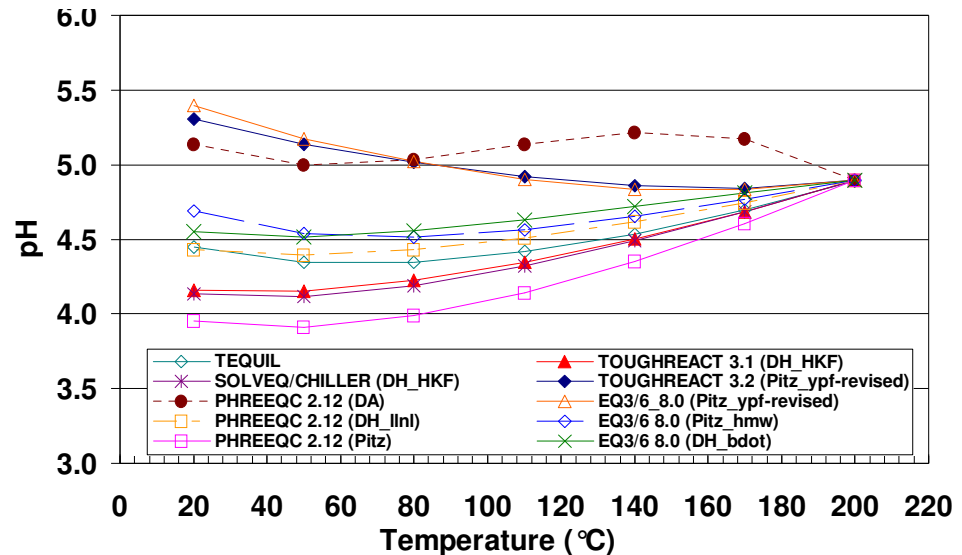
Large uncertainty in A_m and k !!

Beware of Data/Model Limits (!)



Cool same geothermal brine from 200°C (pH 4.9, ionic strength 1.8)

This is mostly an effect of activity coefficient uncertainty at elevated temperature when dealing with concentrated solutions



Useful General References

(not related to TOUGHREACT)

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- Reed M.H., 1982. Calculation of multicomponent equilibria and reaction processes in systems involving minerals, gases, and an aqueous phase. *Geochimica Cosmochimica Acta*, 46, 513-528.
- Steefel, C.I. and Lasaga, A.C. 1994. A coupled model for transport of multiple chemical species and kinetic precipitation/dissolution reactions with application to reactive flow in single phase hydrothermal systems. *American Journal of Science*, 294, (5), 529-592.